

Submesoscale Currents: Surface and Topographic Populations

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Distinctive submesoscale currents are essentially ubiquitous:

1. Near the surface with relatively greater size and strength when horizontal density gradients and/or boundary layer depths are large, with an energy source from mesoscale available potential energy.
2. Near topography where mean or mesoscale currents drag against a sloping bottom in the presence of stratification, generate vertical and potential vorticities, then transiently separate into unstable vortical wakes.

Primary Properties

Scales: horizontal $\sim 10\text{ m} — 10\text{ km}$; vertical $\sim 10\text{ m} — 1\text{ km}$; time $\sim \text{hours} — \text{days}$.
(Note the overlap with IGW scales.)

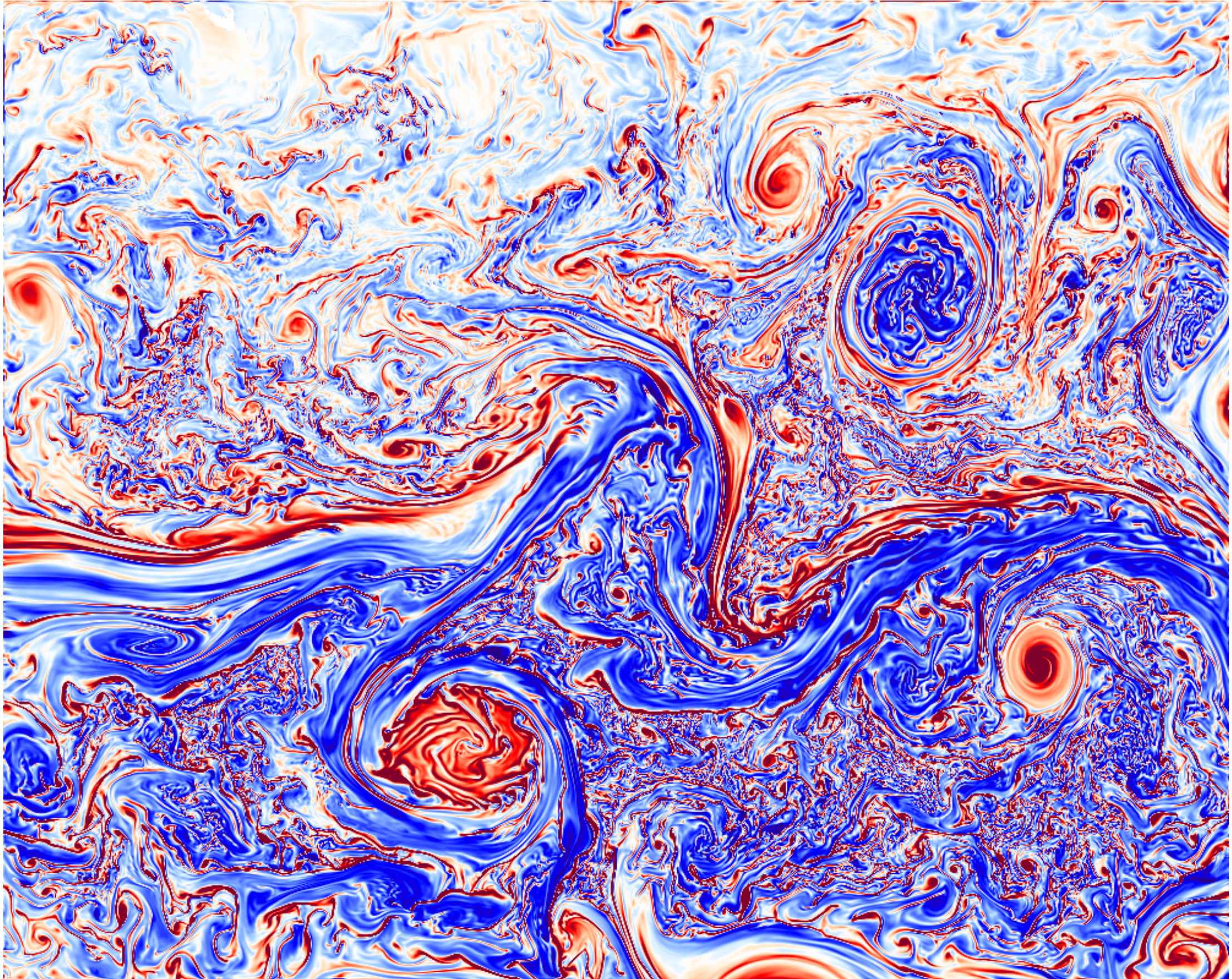
Patterns are surface fronts & dense filaments, topographic wakes, and coherent vortices.

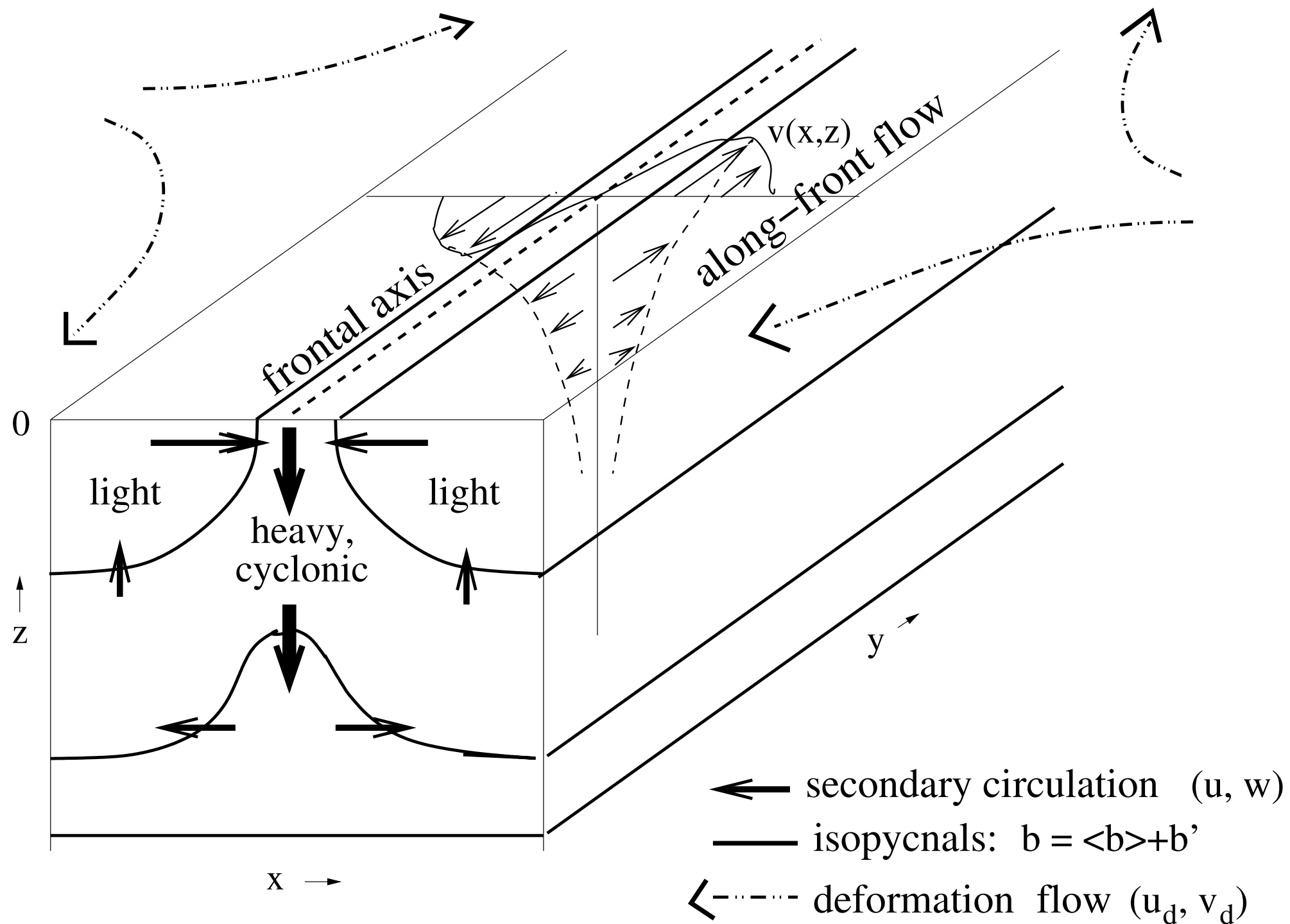
Most evident in **vertical vorticity, vertical velocity, and horizontal density gradient**, in which they stand out from their mesoscale analogs.

Main effects are

- * spontaneous, intrinsic emergence from the mesoscale as a systematic advective transfer of energy to smaller scales;
- * breakdown of geostrophic, hydrostatic balance;
- * lateral material mixing on intermediate lateral scales (or larger in long-lived vortices);
- * vertical material fluxes near the top and bottom boundaries;
- * route to energy dissipation for the general circulation that instigates diapycnal mixing.

Separated Gulf Stream in Winter: surface vorticity

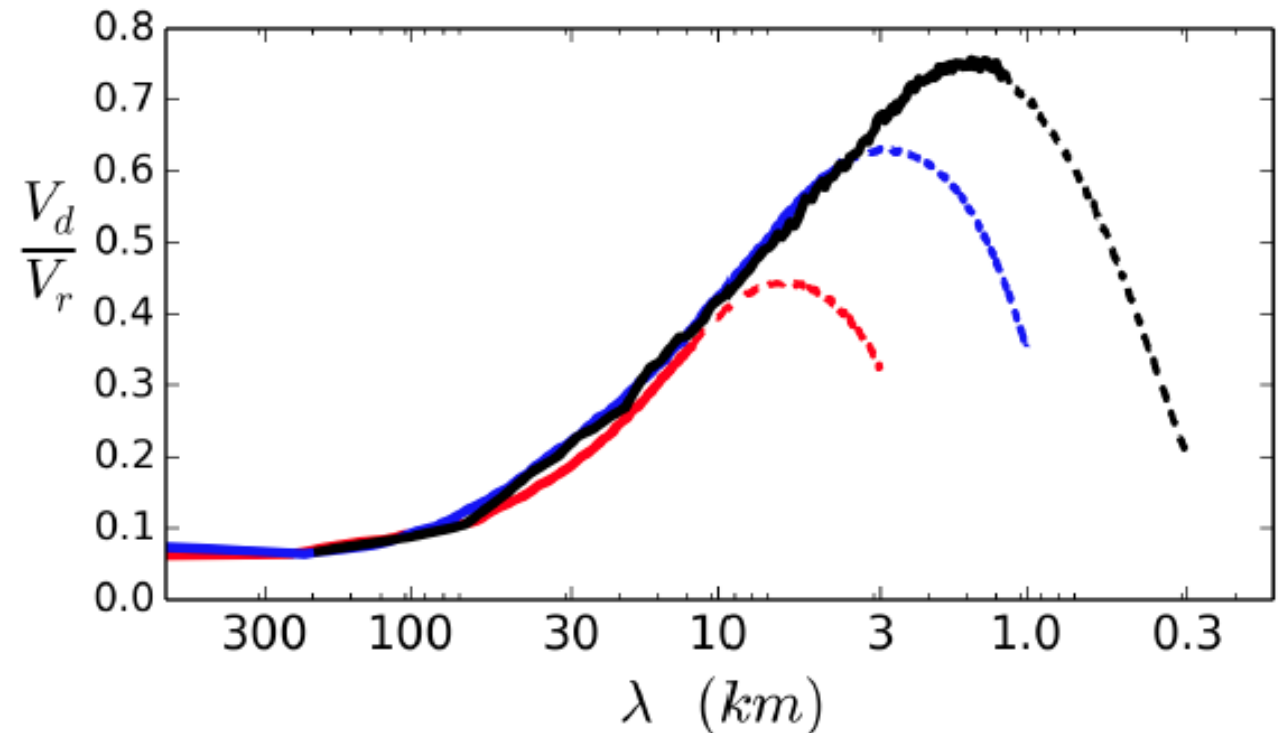
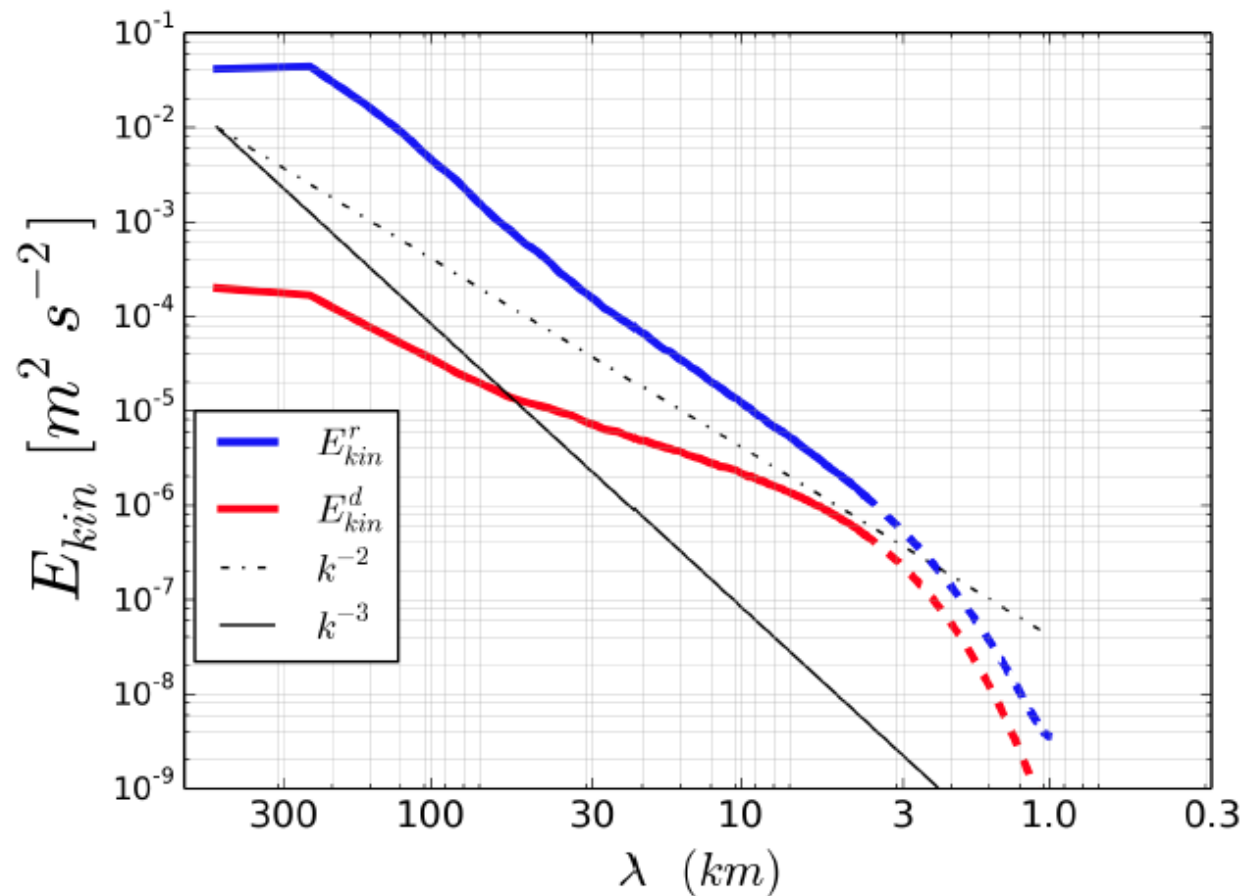




Surface **dense-filament frontogenesis** with secondary circulation and vorticity generation induced by a steady mesoscale deformation (strain) flow. $|\nabla_h \rho|$ and cyclonic vorticity ($f \zeta^z > 0$) grow super-exponentially in time. This process is even stronger than “classical” density-step frontogenesis.

Helmholtz decomposition of horizontal flow at the surface in the Gulf Stream:

$$\mathbf{u}_h = \mathbf{u}_r + \mathbf{u}_d, \quad \nabla \cdot \mathbf{u}_r = 0, \quad \hat{\mathbf{z}} \cdot \nabla \times \mathbf{u}_d = 0.$$



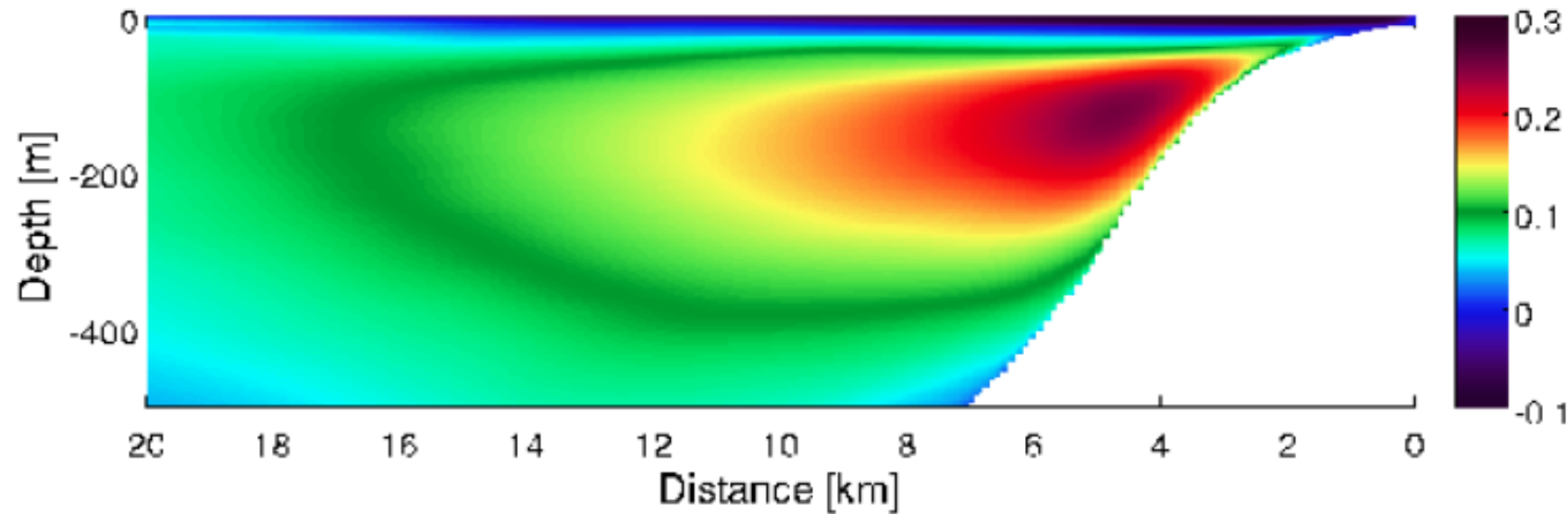
Azimuthally-averaged 2D wavenumber kinetic energy spectra in a simulation with $dx = 0.5$ km. Dashed lines indicate dissipation range.

Square-root ratio of divergent and rotational spectra for 3 solutions with $dx = 1.5, 0.5, \& 0.15$ km. QG theory says this is $O(Ro)$ and must be $\ll 1$.

=> QG behaviors are broken for λ smaller than about 30 km.

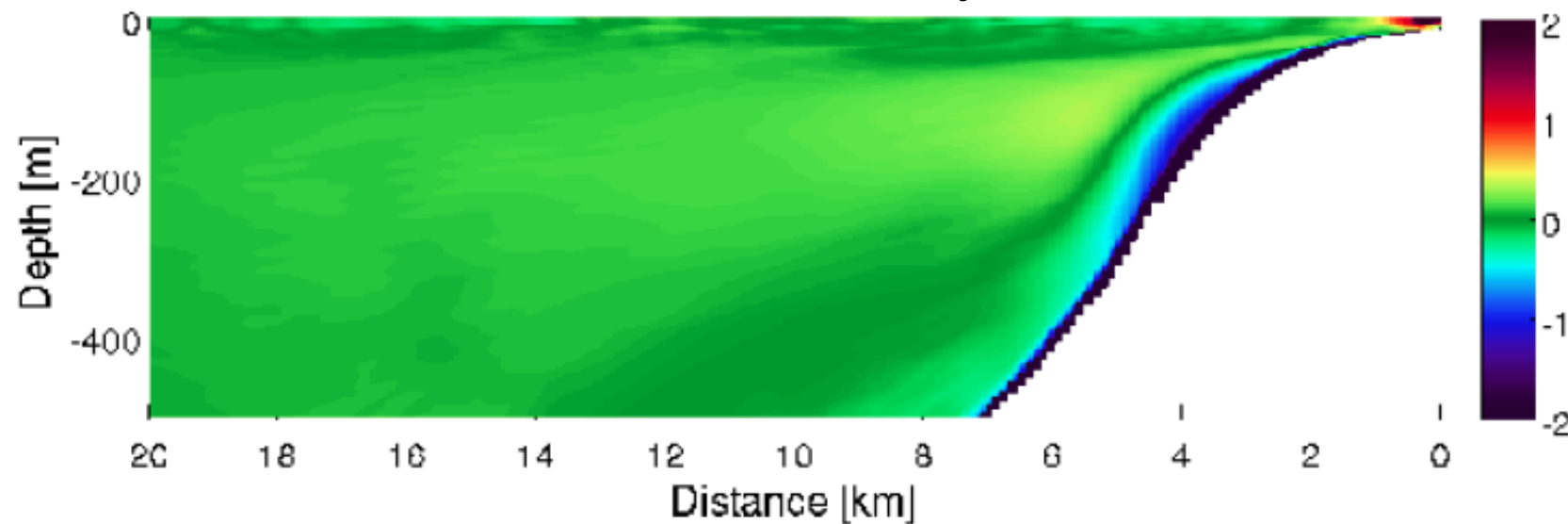
Bottom-boundary Generation of Vertical Vorticity in the California Undercurrent Upstream of a Monterey Bay Separation Point

along-slope velocity [m/s]

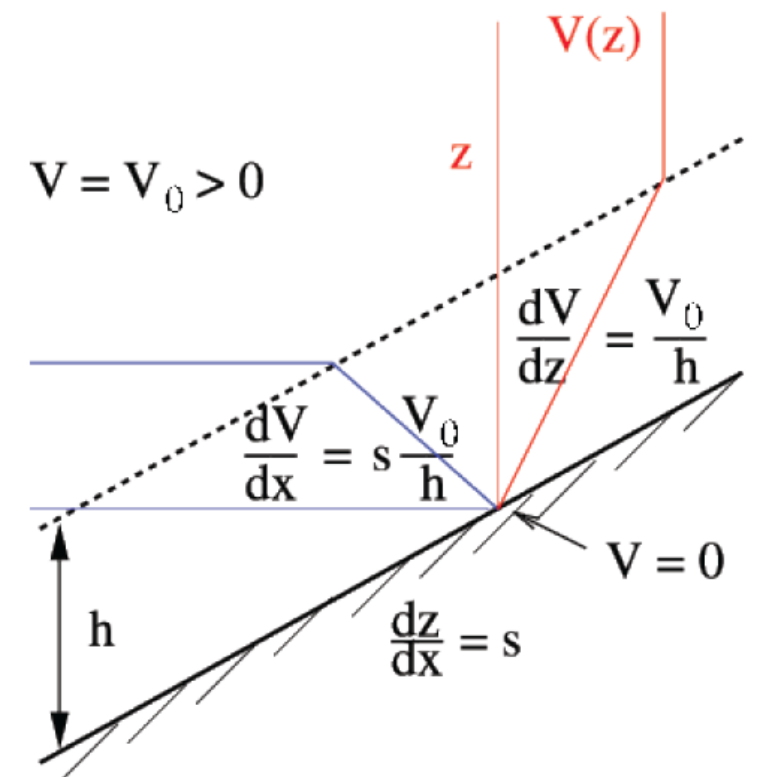


generation mechanism in the turbulent bottom boundary layer:

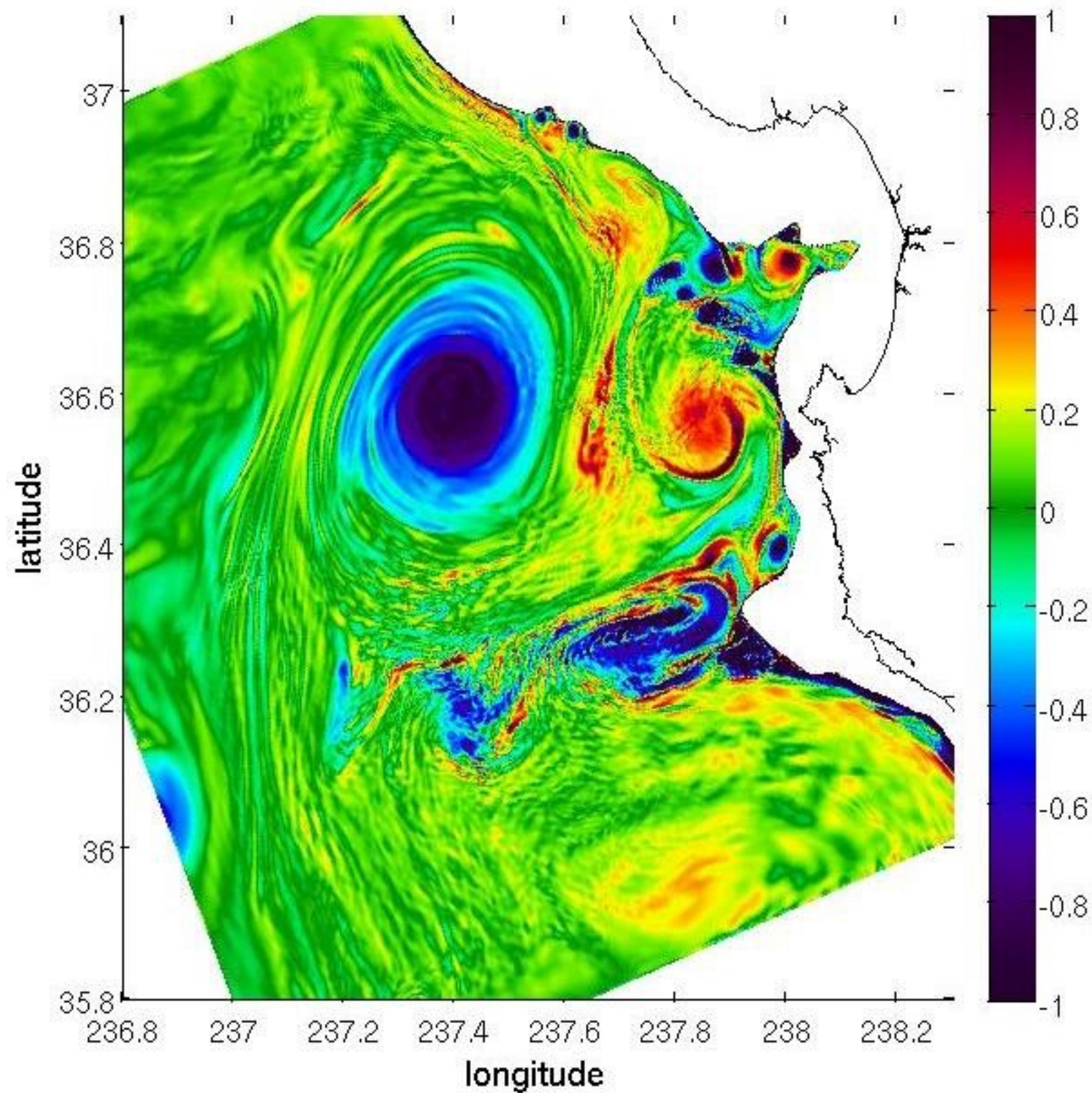
vertical vorticity / f



—> separation, wake instability, and formation of Submesoscale Coherent Vortices (CUDDIES)



(Molemaker et al., 2014)



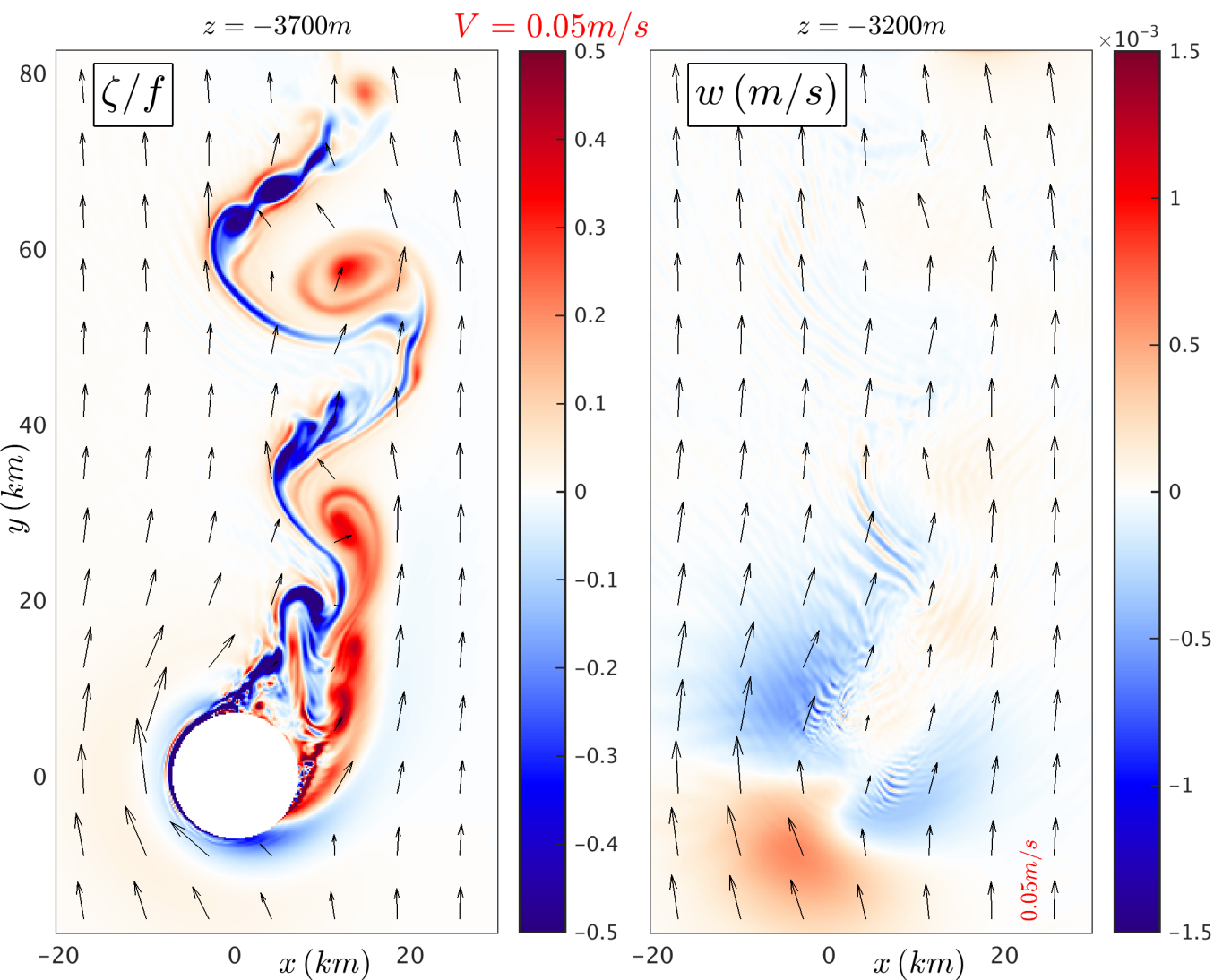
Snapshot of ζ^z / f at 250 m depth in [Monterey Bay](#): the Undercurrent separates at Pt. Sur, goes centrifugally unstable, and eventually reorganizes into “Cuddies”: long-lived Submesoscale Coherent Vortices. Other boundary separation and submesoscale generation sites are also evident.

Uniform inflow past a stratified seamount: Submesoscale-vortical and IGW regimes.

[$H = 4000$ m, $h_{\text{seamount}} = 600$ m, $D_{0.5} = 15$ km, $N/f = 7$]

mid-mount

just above

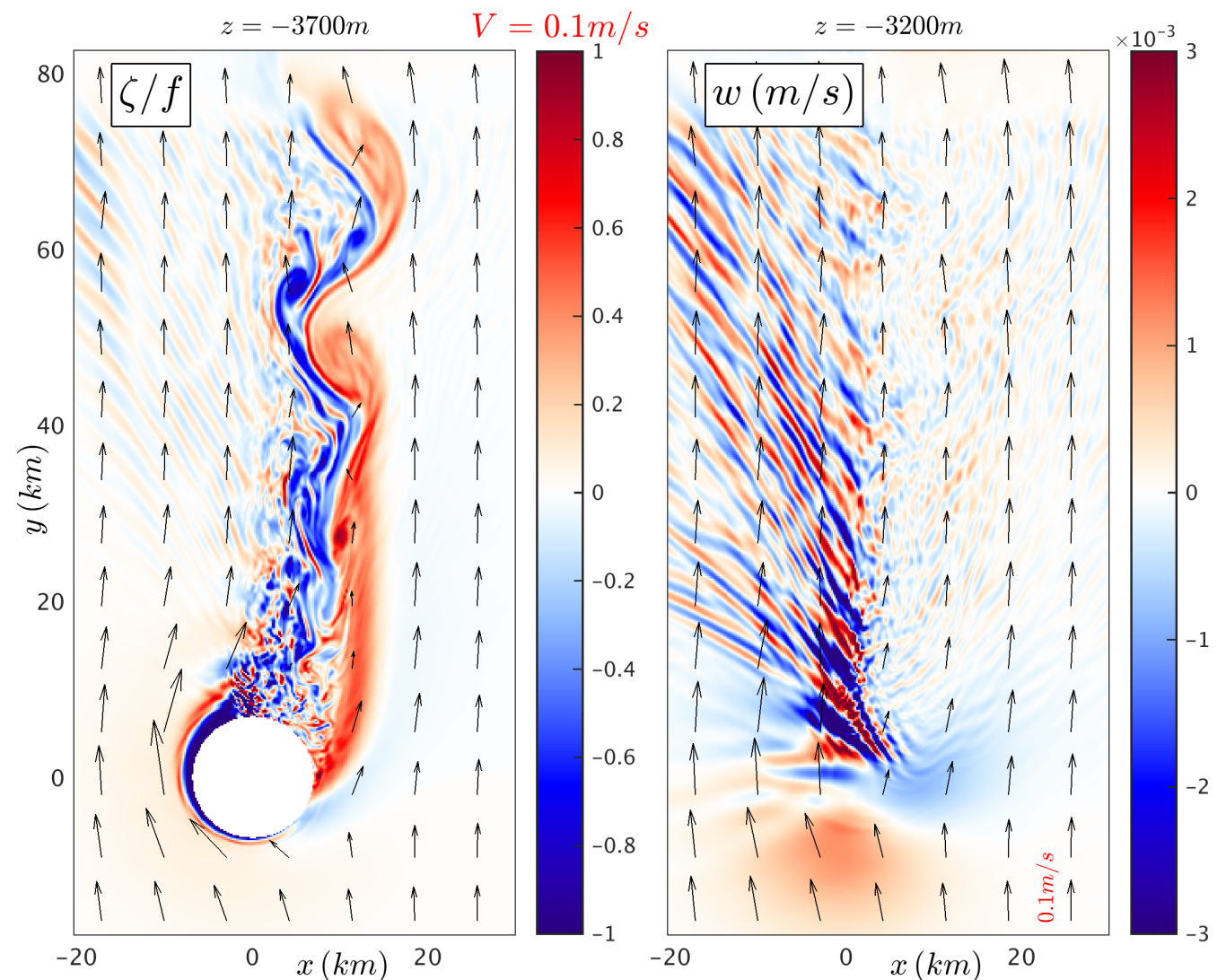


vortical regime

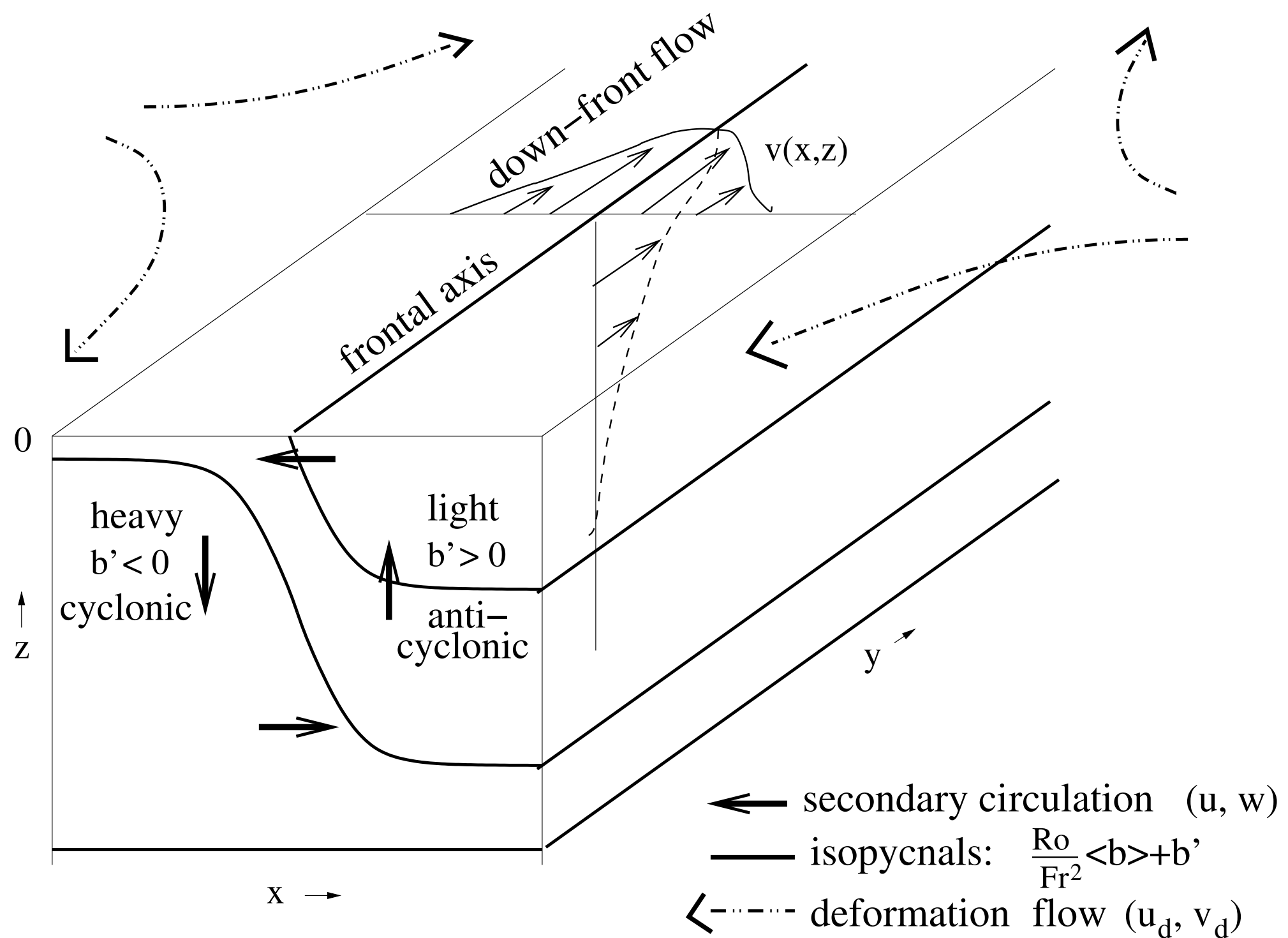
mixed wave & vortical regime

mid-mount

just above

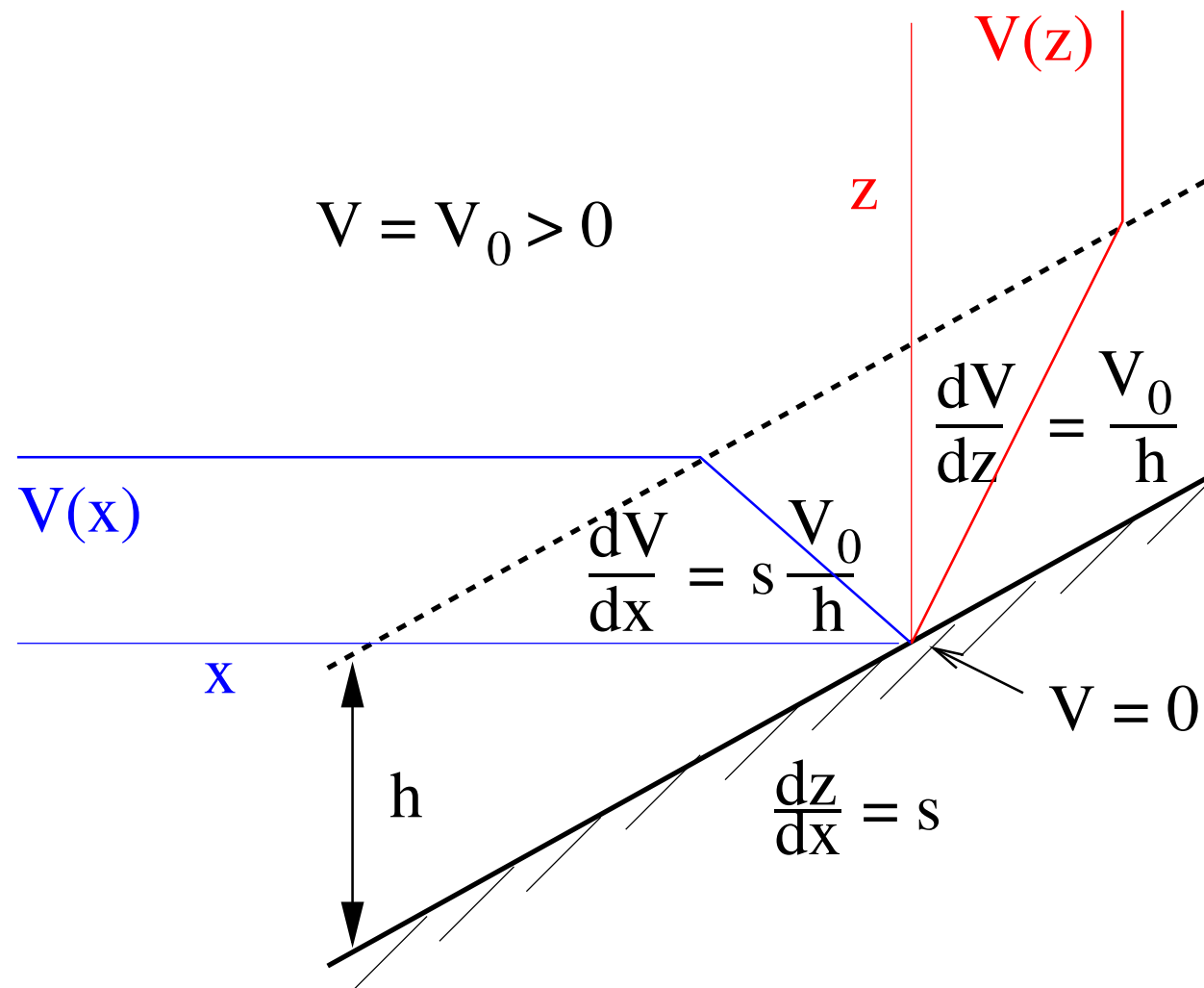


extra slides



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Vertical Vorticity Generation by Along-slope Bottom Flow



Along-slope current V + stratified $b(z)$ + Bottom Boundary Layer (BBL) \Rightarrow

$b(z)$ weakly stratified in bottom layer

$\zeta^z(x) = \partial_x v \sim V s / h$, unlimited by f

flow separation (macro- and mini-wakes)

downstream shear instability (sometimes ageostrophic and centrifugal)

anticyclonic SCV formation by “adjustment” around weakly stratified cores

forward energy cascade \rightarrow high mixing and dissipation in interior